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Weather Derivatives Pricing: Modeling the Seasonal Residual Variance of an Ornstein-Uhlenbeck Temperature Process with Neural Networks

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Abstract

In this paper, we use neural networks in order to model the seasonal component of the residual variance of a mean-reverting Ornstein-Uhlenbeck temperature process, with seasonality in the level and volatility. We also use wavelet analysis to identify the seasonality component in the temperature process as well as in the volatility of the temperature anomalies. Our model is validated on more than 100 years of data collected from Paris, one of the European cities traded at Chicago Mercantile Exchange. Our results show a significant improvement over more traditional alternatives, regarding the statistical properties of the temperature process, which can be used in the context of Monte-Carlo simulations for pricing weather derivatives.

1. Introduction

Since their inception in 1996, weather derivatives have known a substantial growth. Today, weather derivatives are being used for hedging purposes by companies and industries, whose profits can be adversely affected by unseasonal weather or, for speculative purposes by hedge funds and others interested in capitalizing on those volatile markets.

A weather derivative is a financial instrument that has a payoff derived from variables such as temperature, snowfall, humidity and rainfall. However, it is estimated that 98-99% of the weather derivatives now traded are based on temperature. Temperature contracts have as an underlying variable, temperature indices such as Heating Degree Days (HDD) or Cooling Degree Days (CDD) defined on average daily temperatures. The list of traded contracts is extensive and constantly evolving. In the Chicago Mercantile Exchange

(CME) there are traded weather contracts based on an index of Cumulative Average Temperature (CAT) for European cities for May to September. A CAT index is defined as the sum of the daily average temperatures over the period of the contract. Generally, the payoff of a CAT call option at maturity is:

$$c(CAT) = \max(0, (CAT - K))$$

while, the payoff of a CAT put option is:

$$p(CAT) = \max(0, (K - CAT))$$

where K is the strike price of the contract.

However, pricing weather derivatives is far from a straightforward task, since the underlying weather index (HDD, CDD, CAT, etc.) cannot be traded. Furthermore, the corresponding market is relatively illiquid. Consequently, since weather derivatives cannot be cost-efficiently replicated with other weather derivatives, arbitrage pricing cannot directly apply to them. The weather derivatives market is a classic incomplete market, meaning that prices cannot be derived from the no-arbitrage condition, since it is not possible to replicate the payoff of a given contingent claim by a controlled portfolio of the basic securities.

In pricing a weather derivative, dynamic modeling of the daily temperatures is generally considered more appropriate than modeling the temperature index. In principle, it leads to more accurate pricing, but on the other hand deriving an accurate model for the daily temperature is not a straightforward process. Observed temperatures show seasonality in all of the mean, variance, distribution and autocorrelations and long memory in the autocorrelations. The risk with daily

modeling is that small misspecifications in the models can lead to large mispricing in the contracts.

The continuous processes used for modeling daily temperatures usually take a mean-reverting form, which has to be discretized in order to estimate its various parameters. Once the process is estimated, one can then value any contingent claim by taking expectation of the discounted future payoff. Given the complex form of the process and the path-dependent nature of most payoffs, the pricing expression usually does not have closed-form solutions. In that case Monte-Carlo simulations are being used. This approach typically involves generating a large number of simulated scenarios of weather indices to determine the possible payoffs of the weather derivative. The fair price of the derivative is then the average of all simulated payoffs, appropriately discounted for the time-value of money; the precision of the Monte-Carlo approach is dependent on the correct choice of the temperature process and the look back period of available weather data.

In this paper, we address the problem of pricing the European CAT options for the city of Paris. The temperature process on which we build our analysis is the mean-reverting process with seasonality in the level and volatility, proposed by Benth and Saltyte-Benth [1] - a generalisation of the process proposed earlier by Dornier and Querel [2]. This process is discretized in the form of an AR(1) model.

Given the temperature model, the first step of this approach is to identify and remove from the temperature series the (possible) trend and the non-stationary seasonal cycle, hoping that what is left will be stationary. This is usually done by modelling the seasonal variations as deterministic and the same every year (seasonally stationary). The stochastic variability of the temperature is then moved entirely from the seasonal cycle into the residuals.

In modelling the seasonal cycle deterministically, there are three basic approaches: *a*) the averaging method, *b*) the discrete Fourier transform (DFT) and *c*) the regression method. The first approach, calculates average daily temperatures for the year and then it smoothes them. It is the simplest approach, but it is also considered the most inaccurate. In the second approach, the power spectrum of the variance process is estimated, the peaks are reduced to the level of the background and then the power spectrum is adjusted and inverted back to real time. In the third approach, the temperatures are regressed on harmonics of 365 (or 365.25) days. The DFT requires $4N$ years of data, while the regression method can be applied in any number of years. Both methods however, can be used to remove the seasonal cycle both in the mean and in the variance. For a detailed discussion on this subject see Jewson and Brix [3].

Our approach in modelling the seasonal cycle is an extension and combination of the DFT approach and the regression method. More specifically, we use wavelet analysis (WT), -an extension of the DFT which superimposes sines and cosines to represent other functions, to decompose the temperature series into a series of (orthogonal) basis functions (wavelets) with different time and frequency locations. As a result, the wavelet decomposition brings out the structure of the underlying temperature series as well as trends, periodicities, singularities or jumps that could not be observed originally [4], [5]. The information from the wavelet analysis of more than 100 years of temperature data collected from Paris (from 1900 to 2000), is then used in order to identify the trend and select the specific terms of the regression model - a truncated Fourier series.

Once the trend and the seasonal cycle in the mean and the variance have been removed, one has to investigate the distributional properties of the residuals (anomalies) of the temperature process. To the extent that this part of the modeling approach and the initial temperature process are accurate, the residuals must follow a normal distribution with mean zero and standard deviation of one at all times of the year.

However, we find that for the original Ornstein-Uhlenbeck temperature process the hypothesis of normality for the residuals has to be rejected. And the same is true for various extensions of the original model, namely ARMA, ARFIMA and ARFIMA-FIGARCH. This is not surprising, since the temperature for Paris (as for many other locations for which weather derivatives are traded) is non-normally distributed and the above models are Gaussian.

In that case, an approach that could be used is to first transform the temperatures so that they become as close as possible to normal, and then fit the temperature process. However, the disadvantage is that the fitted model does not maximize the likelihood of the original data.

In this paper, we present a novel approach in which we model non-parametrically the variance of the residuals of the Ornstein-Uhlenbeck model with a neural network. Since, we do not make any assumptions regarding the distributional properties of the residuals, this approach is well fitted to deal with difficult problems, like this one, where the distribution of the temperature is not normal.

Since, there is time dependency in the variance of the residuals of the original model, first we extract that variance by grouping the residuals in 365 groups (each group corresponding to a particular day of the year), comprising 101 observations each (number of years in the data set) and then by taking the average for each group. Then, using those 365 values as our data set, we

model the residual variance with a neural network having as inputs the harmonics corresponding to the seasonal cycles of the residuals, identified by a second wavelet analysis.

The improvement regarding the distributional properties of the original model, is significant. The optical examination of the corresponding Q-Q plot reveals that the distribution is quite close to Gaussian, while the Jarque-Bera statistic of the original model is almost halved. Moreover, the observed autocorrelation of the residuals of the original model and that of the residuals of the model after modelling the residual variance with the neural network are quite close. In summary, our approach gives a good fit for the ACF and a reasonable (although quite improved) fit for the residuals.

The rest of the paper is organized as follows. In section 2, we describe the process used to model the average daily temperature in Paris. In section 3, we calibrate the temperature model for Paris based on the results of the wavelet analysis. In section 3.1, we perform wavelet analysis of the temperature series. In section 3.2, we estimate and then remove from the temperature the linear trend. In section 3.3, based on the results of the wavelet analysis we model the seasonality component, we estimate it and then we remove it from the temperature. In section 3.4, we model the seasonal residual variance, again using wavelet analysis as a guide in forming the corresponding model. In section 3.5, we estimate a number of alternative models to the original process, in order to address the observed deviations from normality. In section 4, we present the effect on the temperature process of modelling the residual variance a neural network. Finally, in section 5 we conclude.

2. Dynamic Modeling of the Temperature Process

Many different models have been proposed in order to describe the dynamics of a temperature process. The common assumptions in all these models concerning the temperature are the following: it follows a predicted cycle, it moves around a seasonal mean, it is affected by global warming, it appears to have autoregressive changes and its volatility is higher in winter than in summer.

Early models were using AR(1) processes or continuous equivalents (see for [4], [5], [6]). Other researchers (e.g., [2], [7]) have suggested versions of a more general ARMA(p, q) model. However it has been shown, that all these models fail to capture the slow time decay of the autocorrelations of temperature and hence lead to significant underpricing of weather options [8]. In order to deal with this problem, more

complex models were proposed, with a characteristic example being the model of Brody *et al* [9], which is an Ornstein-Uhlenbeck process. This model was further extended, at first by replacing the noise part of the process (Brownian) by a fractional Brownian noise and then by a Levy process [10].

Our analysis is based on the model of Benth and Saltyte-Benth, where the temperature is expressed as a mean reverting Ornstein-Uhlenbeck process, i.e.

$$dT(t) = dS(t) - \kappa(T(t) - S(t))dt + \sigma(t)dB(t) \quad (1)$$

where, $T(t)$ is the daily average temperature, $B(t)$ is a standard Brownian motion, $S(t)$ is a deterministic function modelling the trend and seasonality of the average temperature, while $\sigma(t)$ is the daily volatility of temperature variations. In [1] both $S(t)$ and $\sigma^2(t)$ are being modeled as a truncated Fourier series, i.e.:

$$S(t) = a + bt + a_0 + \sum_{i=1}^{I_1} a_i \sin(2i\pi(t - f_i)/365) + \sum_{j=1}^{J_1} b_j \cos(2j\pi(t - g_j)/365) \quad (2)$$

$$\sigma^2(t) = c + \sum_{i=1}^{I_2} c_i \sin(2i\pi t / 365) + \sum_{j=1}^{J_2} d_j \cos(2j\pi t / 365) \quad (3)$$

From the Ito formula an explicit solution for (1) can be derived:

$$T(t) = s(t) + (T(t-1) - s(t-1))e^{-\kappa t} + \int_{t-1}^t \sigma(u)e^{-\kappa(t-u)}dB(u) \quad (4)$$

According to this representation $T(t)$ is normally distributed at t and it is reverting to a mean defined by $S(t)$. In this paper, the exact specification of models (2) and (3) is decided based on the results of wavelet analysis of the temperature series.

3. Calibration of the Temperature Model

In this section we derive the characteristics and dynamics of the daily temperature of the city of Paris, France. The data consists of 36,865 values, corresponding to the average daily temperatures of 101 years (1900-2000). In *Figure 1*, we can see the descriptive statistics for the data. The distribution is clearly not

normal, indicating a temperature process that generally hard to model.

In order to identify the number of terms I_1, J_1 in (2) and I_2, J_2 in (3) we decompose the temperature series using a wavelet transform (WT), a generalization of the DFT and the windowed Fourier (WFT) transform. The wavelet transform is localized in both time and frequency. Also it adapts itself to capture features across a wide range of frequencies, thus avoiding the assumption of stationarity. In addition, wavelets have the ability to decompose a signal or a time-series in different levels.

At each level j , we build the j -level approximation a_j , or *approximation* at level j , and a deviation signal called the j -level detail d_j , or *detail* at level j . We can consider the original signal as the approximation at level 0, denoted by a_0 . The words approximation and detail are justified by the fact that a_1 is an approximation of a_0 taking into account the low frequencies of a_0 , whereas the detail d_1 corresponds to the high frequency correction. For detailed expositions on the mathematical aspects of wavelets we refer to (see for example [11], [12], [13]).

3.1. Wavelet Analysis of the Temperature Series in Paris

For the decomposition of the average daily temperature time-series the Daubechies 11 wavelet at level 11 was used. In *Figures 2* and *3*, we can see all the approximations and details (respectively) of the decomposed time-series.

It becomes clear from observing the first seven approximations (a_1 to a_7) and the detail d_8 that there exists a cycle with a period of one year, as it was expected. Approximation a_{11} captures a long cycle with a period of 13 years. Also, in the same approximation an upward trend is observed through the whole period. Detail d_8 also captures a product of two sinusoids, with a period of 1 and 7 years respectively.

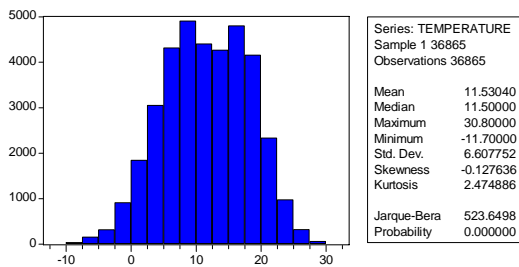


Figure 1. Daily average temperature data distribution statistics for Paris, France for the period 1900-2000.

Details d_{10} and d_{11} reflect a 4-year and an 8-year seasonal effect, respectively. As we can see, both effects are intensive between $t = 1-8,000$ and $t = 20,000-36,865$, while the effects between $t = 8,000-20,000$ are weak. Detail d_9 represents a cycle with period close to 2 years. The visible upward slope, which appears at approximations a_8-a_{11} , reflects the upward trend. The results of wavelet analysis indicate that an upward trend exists throughout the whole period. Finally, the lower details (d_1 and d_2) reflect the noise part of the time-series. A closer inspection of the noise part reveals seasonalities, which will be extracted later on.

3.2. Estimating the Linear Component

A discrete approximation to (4), which is the solution to the mean reverting Ornstein-Uhlenbeck process (1), is:

$$\begin{aligned} T(t+1) - T(t) &= S(t+1) - S(t) \\ &\quad - (1 - e^{-k}) \{T(t) - S(t)\} \\ &\quad + \sigma(t) \{B(t+1) - B(t)\} \end{aligned} \quad (5)$$

which can be written as:

$$\tilde{T}(t+1) = a\tilde{T}(t) + \tilde{\sigma}(t)\varepsilon(t) \quad (6)$$

where

$$\tilde{T}(t) = T(t) - S(t) \quad (7)$$

$$\tilde{\sigma}(t) = a\sigma(t) \quad (8)$$

$$a = e^{-k} \quad (9)$$

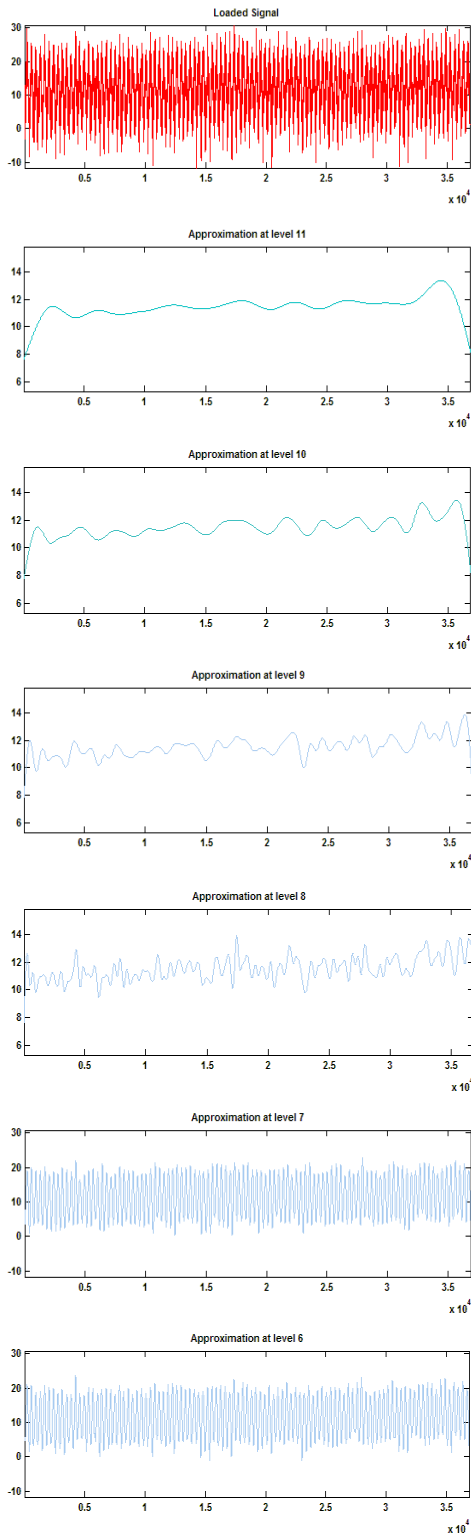


Figure 2. Daily temperature time-series (s) for Paris, France, approximations (a) produced by the wavelet decomposition.

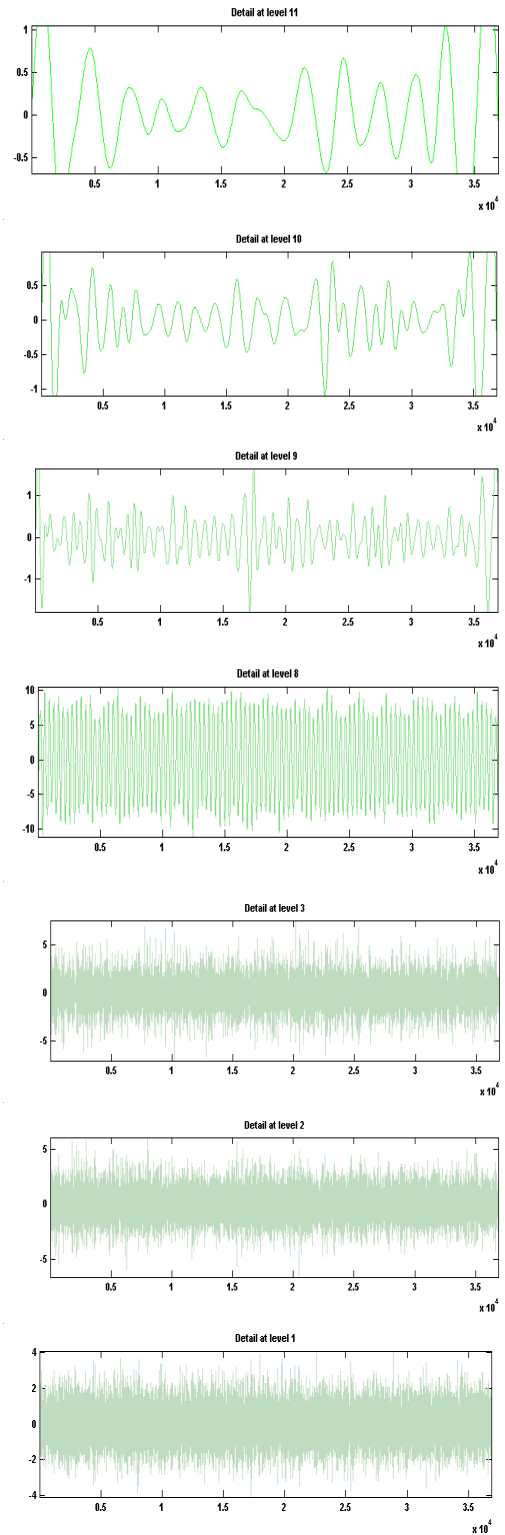


Figure 3. Daily temperature time-series (s) for Paris, France, details (d) produced by the wavelet decomposition.

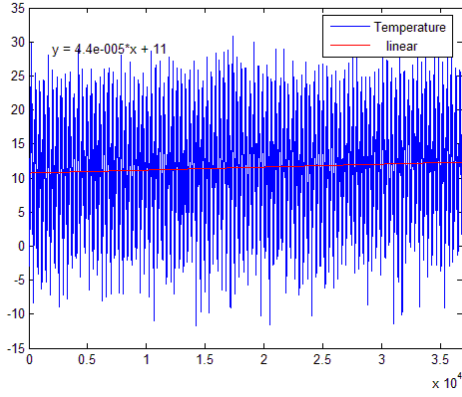


Figure 4. Fitting a trend to the average daily temperature data in Paris for the period from 1900 to 2000.

In order to estimate (6) we need first to remove the trend and seasonality components from the average temperature series.

Firstly, we quantify the upward trend indicated by the results of the wavelet analysis by fitting a linear regression to the temperature data. The regression is statistically significant with intercept $4.3798 \cdot 10^{-5}$ and slope 10.723. The upward trend is depicted in *Figure 4*. Subtracting the trend from the original data we obtain the de-trended temperature series.

3.3. Estimating the Seasonal Component

The results of the wavelet analysis (section 3.1) indicate that the seasonal part of the temperature takes the following form:

$$\begin{aligned}
 S(t) = & a + b_1 \sin(2\pi(t - f_1)/365) \\
 & + b_2 \sin(2\pi(t - f_2)/(2 \cdot 365)) \\
 & + b_3 \sin(2\pi(t - f_3)/(13 \cdot 365)) \\
 & + b_4(1 + \sin(2\pi(t - f_4) \\
 & / (7 \cdot 365))) \sin(2\pi t / 365) \\
 & + b_5 \sin(2\pi(t - f_5)/(8 \cdot 365)) \\
 & + b_6 \sin(2\pi(t - f_6)/(4 \cdot 365))
 \end{aligned} \tag{10}$$

The estimated parameters of the above model are as follows: $a = -0.0001$, $b_1 = -8.0214$, $b_2 = -0.1459$, $b_3 = -0.1421$, $b_4 = 0.1741$, $b_5 = 0.2262$, $b_6 = -0.0223$, $f_1 = -71.4571$, $f_2 = 78.09$, $f_3 = -166.1663$, $f_4 = 787.586$, $f_5 = 598.1549$ and $f_6 = 64.5991$. The mean of the residuals is $-1.5887e-008$ and the standard deviation is 3.4153.

In *Figure 5* the seasonality of the temperature series for the first 10 years, is clearly visible.

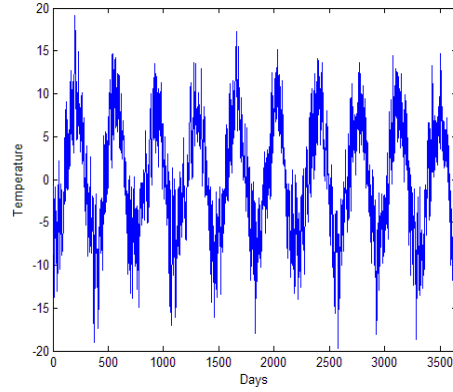


Figure 5. Average daily temperature data in Paris (first 10 years of the time series).

The seasonal component $S(t)$, given by (10), can be seen in *Figure 6*. Next the temperature series is de-seasonalized by removing $S(t)$.

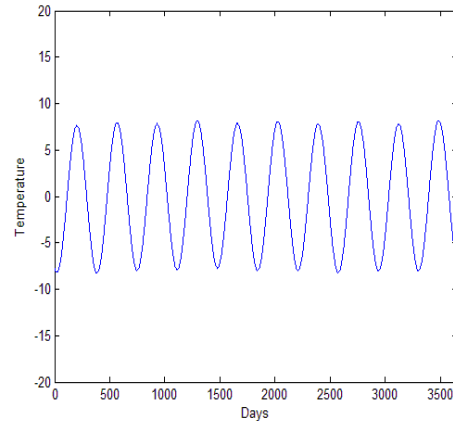


Figure 6. Seasonal component, $S(t)$, of the average daily temperature data in Paris (first 10 years of the time series).

Using the de-trended and de-seasonalized temperature series we estimate the parameters of (6), which is an AR(1) process with zero constant. The mean reversion parameter $a = 0.7978$ and it is statistically significant ($t = 254.05$). The constant is very close to zero, as expected (-0.000989 , $t = 0.01865$). For the overall model, the adjusted $R^2 = 0.6364$ and $F = 64542.19$. In the original continuous-time dynamics model (1), the above value of $a = 0.7978$ corresponds to $k = 0.2259$.

3.4. Modeling the Seasonal Residual Variance

The distributional statistics of the residuals of the AR(1) model (6), indicate a significant deviation from the normal distribution. There is negative skewness (-0.024913), positive kurtosis (3.277200) and the value of the Jarque-Bera statistic is 121.8394. Moreover, the autocorrelation of the residuals is significant for the several first lags (Figure 7), while the autocorrelation of the squared residuals indicates a time dependency in the variance of the residuals (Figure 8). In Figure 8, we can clearly observe a seasonal variation.

Since, for the residuals $e(t)$ of the AR(1) is true that

$$e(t) = \tilde{\sigma}^2(t)\varepsilon(t) \quad (11)$$

where $\varepsilon(t)$ are *i.i.d.* $N(0,1)$, we can extract the variance $\tilde{\sigma}^2(t)$ as follows: Firstly, we group the residuals in 365 groups, comprising 101 observations each (each group corresponds to a single day of the year). Then, by taking the average of the squares of each group we obtain the variance.

From (8) it is true that:

$$\sigma^2(t) = \frac{\tilde{\sigma}^2(t)}{a^2} \quad (12)$$

where $a = 0.7978$.

In deciding which terms of a truncated Fourier series to use in order to model the variance $\sigma^2(t)$ (its empirical values are being computed using equation 12), we perform again a wavelet analysis, which indicates the presence of five cycles within $\sigma^2(t)$. A one-year cycle, a half-year cycle, a 1/4 of a year cycle, a 1/9 of a year cycle and a 1/18 of a year cycle. We model accordingly the variance $\sigma^2(t)$, as follows:

$$\begin{aligned} \sigma^2(t) = & c_0 + c_1 \sin(2\pi t / 365) \\ & + c_2 \sin(4\pi t / 365) \\ & + c_3 \sin(8\pi t / 365) \\ & + c_4 \sin(18\pi t / 365) \\ & + c_5 \sin(36\pi t / 365) \\ & + d_1 \cos(2\pi t / 365) \\ & + d_2 \cos(4\pi t / 365) \\ & + d_3 \cos(8\pi t / 365) \\ & + d_4 \cos(18\pi t / 365) \\ & + d_5 \cos(36\pi t / 365) \end{aligned} \quad (13)$$

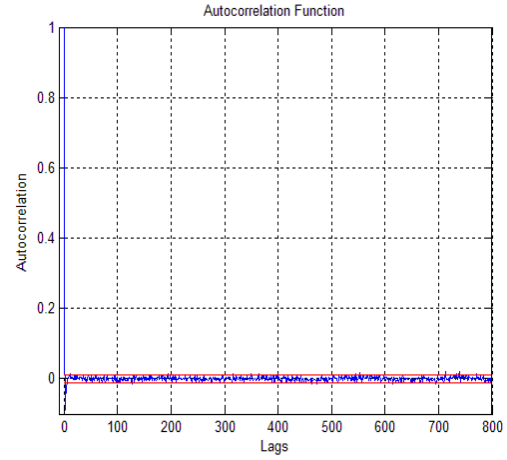


Figure 7. ACF for the residuals of the AR(1) model of the de-trended and de-seasonalized Paris average daily data.

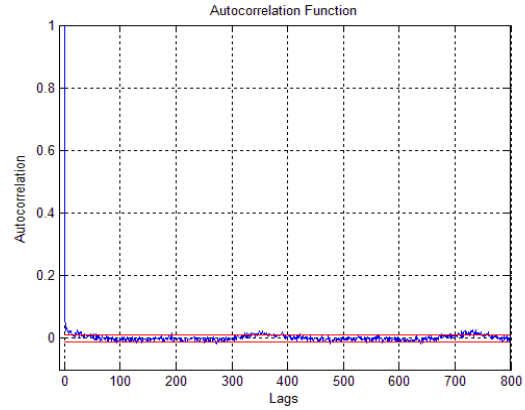


Figure 8. ACF of the squared residuals of the AR(1) model of the de-trended and de-seasonalized Paris average daily data.

The values of the estimated parameters of (13) are: $c_0 = 4.2398$, $c_1 = 0.4324$, $c_2 = -0.2641$, $c_3 = 0.0557$, $c_4 = 0.0843$, $c_5 = -0.0131$, $d_1 = 0.5610$, $d_2 = 0.6195$, $d_3 = 0.0326$, $d_4 = 0.0161$ and $d_5 = -0.0421$.

The empirical values of the variance of the residuals (365 values) together with the fitted variance

$$\tilde{\sigma}^2(t) = a^2 \sigma^2(t) \quad (14)$$

can be seen in Figure 9. We observe that the variance takes its highest values during the winter months, while it takes its lowest values during early Autumn.

The standard deviation of the residuals is 0.6035, while the standard deviation of the remaining noise part is 1.0003 and its mean is 0.0018.

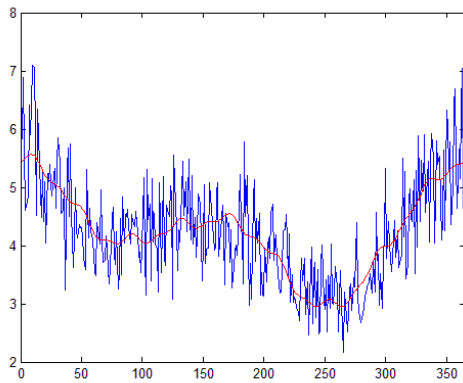


Figure 9. Empirical variance and fitted variance $\tilde{\sigma}^2(t)$.

In Figure 10, we can see the autocorrelation function of the squared residuals of the AR(1) process after dividing out the volatility (14) from the regression residuals.

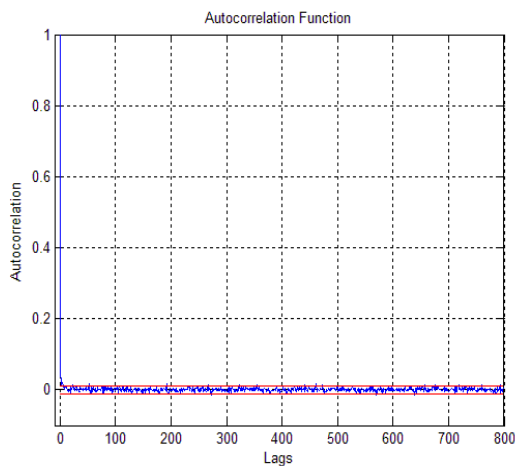


Figure 10. ACF of the squared residuals of the AR(1) model after dividing out the volatility function $\tilde{\sigma}(t)$ from the regression residuals.

We observe that the seasonality has been removed, but there is still autocorrelation in the first three lags. Moreover, since the Jarque-Bera statistic is 67.6 with a p -value of 0.000000, we have to reject the hypothesis of normal distribution.

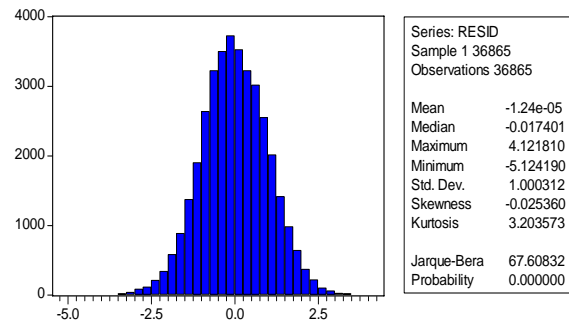


Figure 11. Distribution statistics of the residuals of the AR(1) model after dividing out the volatility function $\tilde{\sigma}(t)$ from the regression residuals.

3.5. Dealing with Non-Normality

The findings of Benth and Saltyte-Benth [10] for the Copenhagen temperature series are very similar. Although, they did not use wavelet analysis to calibrate their models, they had managed to remove seasonality from the residuals, but their distribution proved to be non-normal. They suggested that a more refined model would probably rectify this problem, but they did not proceed in estimating one. In an earlier paper regarding Norwegian temperature data, Benth and Saltyte-Benth [1] suggested to model the residuals by a generalized hyperbolic distribution. However, as the same authors comment the inclusion of a non-normal model leads to complicated Levy process dynamics.

We estimated a number of alternatives to the original AR(1) model. In particular we estimated an ARMA(3,1) model, a long-memory homoscedastic ARFIMA model and a long-memory heteroscedastic ARFIMA-FIGARCH model.

As we can see in the respective Q-Q plots of the residuals in Figures 12, 13 and 14 the hypothesis of normality has to be rejected. The Jarque-Bera statistic is above 91 for the ARMA(3,1) model, above 94 for the ARFIMA model and above 114 for the ARFIMA-FIGARCH model. In contrast with [8] and [14] our results represent a disimprovement of our original AR(1) process. As the complexity of the models increases, the Jarque-Bera statistic becomes larger which suggest that the AR(1) model is the most appropriate.

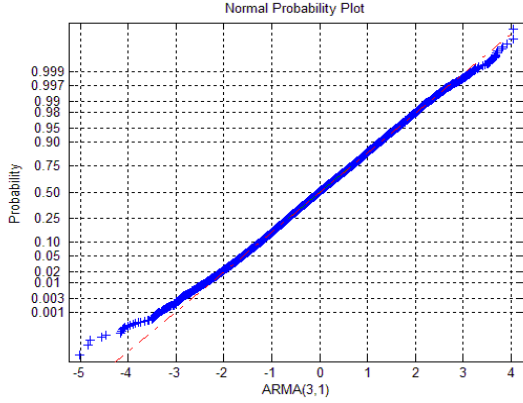


Figure 12. Q-Q plot of the residuals of an AR-MA(3,1) model.

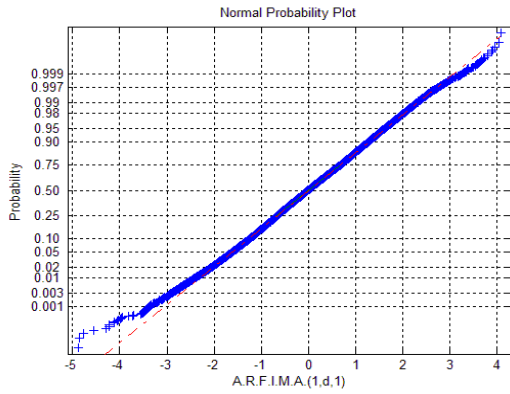


Figure 13. Q-Q plot of the residuals of an AR-FIMA model.

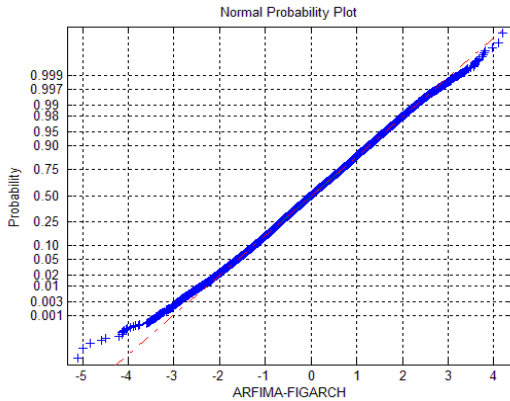


Figure 14. Q-Q plot of the residuals of an AR-FIMA-FIGARCH model.

Concluding, although the AR(1) model probably is not the best model for describing temperature anomalies, increasing the model complexity (ARMA, ARFIMA, ARFIMA-FIGARCH) and thus the complexity of theoretical derivations in the context of weather derivative pricing does not seem to be justified.

4. Modeling the Seasonal Residual Variance with Neural Networks

Next, we model the seasonal residual variance with a backpropagation neural network, with one hidden layer, a bias term and ten hidden units (Figure 15). For a single-hidden-layer architecture, the number of hidden units λ is an unambiguous descriptor of the dimensionality p of the parameter vector; $p = (m + 2)\lambda + 1$. In this case $p = (10 + 2)10 + 1 = 121$. The corresponding number of observations/parameters ratio is relatively low ($n/p \approx 3$).

4.1. The Neural Network Model

Our hypothesis here is that between the seasonal variance and the ten harmonics identified by the wavelet analysis there is a deterministic relationship $\varphi(\bullet)$ of the general form:

$$\sigma^2(t) = \varphi(\sin(2\pi t / 365), \sin(4\pi t / 365), \sin(8\pi t / 365), \sin(18\pi t / 365), \sin(36\pi t / 365), \cos(2\pi t / 365), \cos(4\pi t / 365), \cos(8\pi t / 365), \cos(18\pi t / 365), \cos(36\pi t / 365)) \quad (15)$$

We estimate $\varphi(\bullet)$ non-parametrically with the neural network $g(\bullet)$. Given an input vector \mathbf{x} (the harmonics) and a set of weights \mathbf{w} (parameters), the network response (output) $g_\lambda(\mathbf{x}; \mathbf{w})$ is:

$$g_\lambda(\mathbf{x}; \mathbf{w}) = \gamma \left(\sum_{j=1}^{\lambda} w_j^{[2]} \gamma \left(\sum_{i=1}^m w_{ij}^{[1]} x_i + w_{m+1,j}^{[1]} \right) + w_{\lambda+1}^{[2]} \right) \quad (16)$$

where, $w_{i,j}^{[1]}$ is a weight corresponding to the connection between the i^{th} input and the j^{th} hidden unit, $w_{m+1,j}^{[1]}$ is a bias term corresponding to the j^{th} hidden unit, $w_{\lambda+1}^{[2]}$ is the weight of the connection between the j^{th} hidden unit and the output unit, and $w_{\lambda+1}^{[2]}$ is the bias term of the output unit, and the function $\gamma(\bullet)$ is a sigmoidal function.

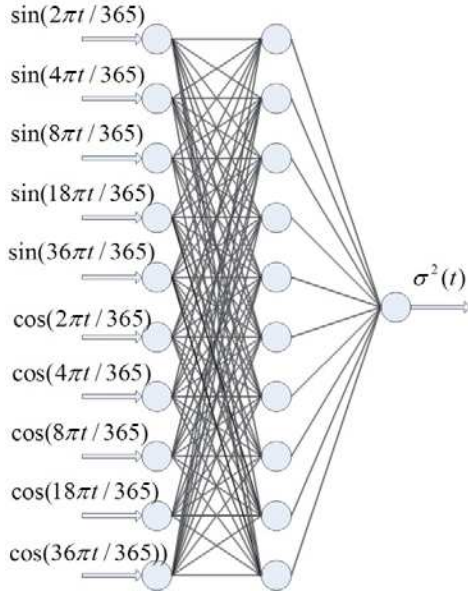


Figure 15. A neural network for modelling the seasonal residual variance.

An estimate of the parameter vector \mathbf{w} is obtained by minimising iteratively the ordinary least squares cost functional:

$$L_n(\mathbf{w}) = \frac{1}{2n} \sum_{i=1}^n (\sigma^2(t) - g(\mathbf{x}_i; \mathbf{w}))^2 \quad (17)$$

where $\sigma^2(t)$ is the t^{th} target output variance and n is the number of training vectors (365). The loss function in (17) gives us a measure of accuracy with which the estimator fits the observed data but it does not account for the estimator's (model) complexity. Given a sufficiently large number of free parameters, p , the neural estimator can fit the data with arbitrary accuracy.

4.2. Removing the Irrelevant Connections

Once, the parameters of the neural model (16) are estimated, we have to deal with the presence of flat minima (potentially many combinations of the network parameters corresponding to the same level of the empirical loss), especially if the statistical properties of the model are of importance, as it is the case in this complex financial application. In order to identify a locally unique solution, we have to remove all the irrelevant parameters, that is the parameters that do not affect the level of the empirical loss.

For this purpose we use the Irrelevant Connection Elimination scheme (ICE) [15], which is much less computationally demanding than other alternatives

since, although it uses the full Hessian of L_n , it does not require inverting the Hessian matrix – a common requirement of other algorithms. ICE is based on the Taylor's approximation of the empirical loss:

$$\begin{aligned} \delta L_n = & \sum_{i=1}^p g_i \delta w_i + \frac{1}{2} \sum_{i=1}^p a_{ii} \delta w_i^2 \\ & + \frac{1}{2} \sum_{i=1, i \neq j}^p a_{ij} \delta w_i \delta w_j + O((\delta w)^3) \end{aligned} \quad (18)$$

From (18) ICE derives the "saliencies" $S(w_i)$, i.e., the contribution of w_i to δL_n , when a small perturbation δw_k is added to all connections. as:

$$S(w_i) = g_i \delta w_i + \frac{1}{2} \sum_{j=1}^p a_{ij} \delta w_i \delta w_j \quad (19)$$

$$\text{where, } \delta L_n = \sum_{i=1}^p S(w_i).$$

At a well-defined local minimum (19) can be simplified by setting $g_i = 0$, although this is not a requirement. The method can be summarised in the following steps:

Step 1: Train to convergence.

Step 2: Compute the saliencies $S(w_i)$

Step 3: Deactivate the connection with the least associated saliency, unless it was reactivated in step 5. When a pre-specified maximum number of steps is been reached, then the algorithm STOPS.

Step 4: Train further for a small number of epochs, until the training error has stabilised.

Step 5: If the training error has increased, reactivate the connection, otherwise remove it. Then go to step 3.

Because of possible dependencies in the connections, it is not advisable to remove more than one connection at a time (the removal of one connection can affect the standard errors and saliencies of others). This does not pose any computational problems to ICE, since computing the Hessian is the same order of complexity as computing the derivatives $\partial L_n / \partial w_i$ during training.

After applying the algorithm ICE to the estimated network, the number of parameters is reduced to 29 out of 121 originally. The observations/parameters ratio now takes the value of $n/p = 12.6$ (from 3 originally). Roughly speaking, the reduced network is the same order of complexity with a fully connected network with two hidden units.

4.3. Statistics for the Neural Model and the Residuals of the Ornstein-Uhlenbeck Process

The summary statistics for the reduced neural model of the seasonal variance (removing the irrelevant connections) are given in *Table 1*. As we can see, the coefficient of determination (R^2) of the model (adjusted for degrees of freedom), as well as the R^2 of the linear regression of the predicted variances vs. the target variances are quite high.

Table 1: Neural Model Summary Statistics

Average Squared Error (ASE)	0.007778
Standard error of the estimate (SE)	0.088193
Mean absolute error (MAE)	0.070264
Empirical loss (L_n)	0.003889
Generalised Cross Validation (GCV)	0.017405
Final Prediction Error (FPE)	0.015492
R -squared	64.63906
R -squared (adjusted for d.f.)	59.22428
R -squared for the linear regression of forecasted variance vs. target variance	64.97905

In *Table 2* we can see the variable significance statistics for the network inputs (the input X_1 corresponds $\sin(2\pi t/365)$, and so on). The relevance S of the input variables to the model is quantified by the partial derivative of the empirical loss (17) to this variable. The sampling variance of the relevance metric is quantified by performing random sampling from the limiting joint distribution of the parameters (it is assumed multivariate normal). For more details on this technique refer to [16].

Table 2: Neural Model Variable Significance Estimation Statistics

VAR	$S=dL_n/dX$	St.Dev.	t -value
X_7	0.00243	0.00020	12.37317
X_1	0.00184	0.00014	13.00622
X_6	0.00184	0.00014	12.89321
X_2	0.00158	0.00316	0.49970
X_{10}	0.00129	0.00259	0.49982
X_3	0.00091	0.00006	15.16516
X_4	0.00076	0.00005	16.57073
X_8	0.00057	0.00114	0.49975
X_9	0.00024	0.00007	3.65349
X_5	0.00006	0.00011	0.50023

Input variables in *Table 2* are sorted in descending magnitude of the relevance metric S . The most impor-

tant and statistically significant variables appear to be X_7, X_1, X_6, X_3, X_4 and X_9 , i.e.,

$$\begin{aligned} &\cos(4\pi t/365) \\ &\sin(2\pi t/365) \\ &\cos(2\pi t/365) \\ &\sin(8\pi t/365) \\ &\sin(18\pi t/365) \\ &\cos(18\pi t/365) \end{aligned}$$

According to this, the one year cycle (the terms containing $2\pi t$) and the $1/9^{\text{th}}$ of a year cycle (the terms containing $18\pi t$) appear to be significant. The half year cycle (the term containing $4\pi t$) and the $1/4^{\text{th}}$ of the year cycle (term containing $8\pi t$) also appear to be significant. We also note that since the input $\cos(2\pi t/365)$ does not appear to be significant, the half year cycle takes its highest value at the beginning of each year. Also since the input $\cos(8\pi t/365)$ does not appear to be significant, the $1/4^{\text{th}}$ of the year cycle takes the value of zero at the beginning of each year.

Proceeding now to the analysis of the residuals of the Ornstein-Uhlenbeck temperature process, when using the neural network for estimating nonparametrically the seasonal variance, the first thing we observe in the ACF of the residuals (*Figure 16*) is that the first three lags are significant. This is in line with the observed behaviour of the temperature series.

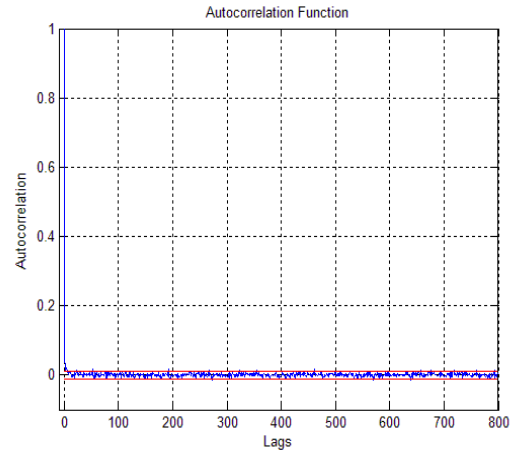


Figure 16. ACF of the squared residuals of the Neural Network after dividing out the volatility function $\tilde{\sigma}(t)$ from the regression residuals.

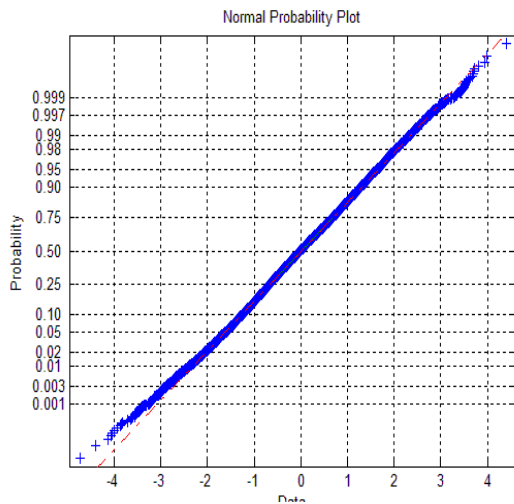


Figure 17. Q-Q plot of the residuals of the Neural Network calculated from 101 years of data.

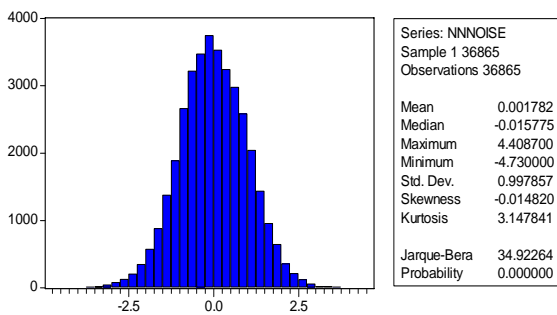


Figure 18. Distributional statistics for the residuals of the Neural Network calculated from 101 years of data.

Furthermore, the improvement regarding the distributional properties of the original model, is significant. The optical examination of the Q-Q plot of the residuals in *Figure 17*, reveals that the distribution is quite close to Gaussian, while the Jarque-Bera statistic of the original model is almost halved (*Figure 18*).

5. Summary

In this paper, in the context of an Ornstein-Uhlenbeck temperature process we have used wavelet analysis to identify the seasonality component in the temperature process as well as in the volatility of the residuals, for the average daily temperature in Paris. The temperature anomalies, however, deviated to some extent from normality. In an attempt to rectify this problem, we estimated a number of alternatives to the

original AR(1) model. In particular we estimated an ARMA(3,1) model, a long-memory homoscedastic ARFIMA model and a long-memory heteroscedastic ARFIMA-FIGARCH model. However, none of these alleviated the problem.

In the next step of our analysis we have used a neural network to model the seasonal volatility of the residuals. We “localized” the neural model, i.e., we removed the irrelevant connections of the estimated model, with the algorithm ICE, which quantifies the contribution of each connection to the change of the empirical loss, when a small perturbation is added to all connections.

Employing a neural network to the estimation of the seasonal volatility of the residuals, led to a significant improvement regarding the distributional properties of the original model.

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